Assessing aeolian beach-surface dynamics using a remote sensing approach

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**Abstract**

A remote sensing technique for assessing beach surface moisture was used to provide insight into beach-surface evolution during an aeolian event. An experiment was carried out on October 21, 2007 at Greenwich Dunes, Prince Edward Island National Park, Canada, during which cameras were mounted on a mast on the foredune crest at a height of about 14 m above the beach. Maps of beach surface moisture were created based on a calibrated relationship between surface brightness from the photographs and surface moisture content measured in situ at points spaced every 2.5 metres along a transect using a Delta-T moisture probe. A time sequence of maps of surface moisture content captured beach surface evolution through the transport event at a spatial and temporal resolution that would be difficult to achieve with other sampling techniques such as impedance probes. Erosion of the foreshore and berm crest resulted in an increase in surface moisture content in these areas as the wetter underlying sediments were exposed. Flow expansion downwind of the berm crest led to sand deposition and a consequent decrease in surface moisture content. Remote sensing systems such as the one presented here allow observations of the combined evolution of beach surface moisture, shoreline position, and fetch distances during short-term experiments and hence provide a comprehensive rendering of sediment erosion and transport processes.

**Keywords:** beach surface moisture, remote sensing, aeolian geomorphology, Greenwich Dunes
1. Introduction

Surface moisture is an important control on the threshold of movement for aeolian transport and likely on the instantaneous transport rate (Belly, 1964; McKenna Neuman and Nickling, 1989; Namikas and Sherman, 1995; Cornelis and Gabriels, 2003; Wiggs et al., 2004a; Davidson-Arnott et al., 2008; Nordstrom et al., 2011). On beaches, surface moisture can be largely variable over time and space, and it is subject to change during aeolian transport events due to local erosion that exposes underlying wetter sediments or local accumulation of dry sand on a previously wetted surface (e.g. Nield et al., 2011). Until recently, measurements of surface moisture were acquired by taking surface scrapings or cores, and returning the samples to the laboratory for gravimetric analysis. In addition to the considerable labour involved and the difficulty of relating the results to the true moisture content of the surface grain layer itself, this technique resulted in physical disturbance of the surface through sampling and trampling. The development of impedance probes such as the Delta-T Theta probe (Atherton et al., 2001; Wiggs et al., 2004b; Yang and Davidson-Arnott, 2005; Edwards and Namikas, 2009) allowed more frequent measurements of moisture content during transport events, but such probes still have the disadvantage that readings need to be taken at many points and that each measurement integrates moisture over a sampling depth of about 2-6 cm depending on the length of the probe.

The development of remote sensing technologies for assessing surface moisture using the apparent surface brightness offers the potential to sample
instantaneously over a substantial area without disturbing the surface, and this approach provides a better measure of the true surface moisture content (McKenna Neuman and Langston, 2006; Darke et al., 2009; Delgado-Fernandez et al., 2009). Remote sensing systems have become increasingly important tools in coastal research, with a number of processes and forms now being monitored using Argus Systems (Holman and Stanley, 2007), satellite imagery (e.g., Wahid, 2008), Terrestrial Laser Scanner (TLS; Nield and Wiggs, 2011) and time-lapse photography (e.g., Lynch et al., 2006; Darke et al., 2009; Delgado-Fernandez et al., 2009). Summerfield (2005, p. 403) notes that it is often the introduction of new techniques that stimulates ‘leaps of explanatory power’ and permits investigation into questions previously discarded as ‘incapable of being addressed’. In particular, the deployment of camera systems for periods of months to years has provided rich insight into the nature of aeolian transport events in coastal areas (Delgado-Fernandez and Davidson-Arnott, 2011).

The primary purpose of this paper is to discuss the use and potential of a remote-sensing camera system (Delgado-Fernandez et al., 2009) as part of a short-term field experiment designed to measure wind flow and sand transport across a beach foreshore and berm over a period of several hours. In this application imagery was captured every 5 minutes, in contrast to the hourly data collected for the meso-scale monitoring (Delgado-Fernandez, 2011). This allowed detailed assessment of the evolution of surface moisture during the transport event with the prospect of evaluating the effect of moisture on the spatial and temporal variations in the sand transport rate. Information derived
from the remotely sensed imagery was compared to data from the Theta probe in
order to determine the accuracy and precision of each methodology. The
discussion in this paper focuses on the effects of aeolian transport on surface
moisture content rather than moisture as a control on the threshold of aeolian
transport. While the latter has been widely discussed, the former can now be
investigated by analysing the complexities shown in moisture maps obtained
from the remote sensing system.

2. Study site

The study site was located at Greenwich Dunes, Prince Edward Island
(PEI), Canada (Figure 1). Greenwich Dunes are part of a barrier spit beach-dune
complex formed on the east side of the entrance to Saint Peter’s Bay, on the
north shore of PEI (Mathew et al., 2010). The coast is microtidal with a mixed
semidiurnal regime and a maximum range at spring tides of approximately 1 m.
The foredune ranges in height from 6 to 10 m with a stoss slope of 20-25° and
the dune crest is aligned roughly east-west. The beach is 30-40 m wide and
consists predominantly of quartz sand with a mean diameter of 0.26 mm. Marram
grass (*Ammophila breviligulata*) covers the foredunes and exhibits considerable
seasonal patterns of density and height, with individual plants typically reaching
25 cm heights. The study site was located approximately 1 km east of Saint
Peter’s Bay, and covered an area of approximately 60 m alongshore. Further
information on the study site may be found in Hesp et al. (2005), Walker et al.
(2006), Davidson-Arnott et al. (2008) and Bauer et al. (2009).
3. Methods

Wind and sand transport intensity

Wind and sand transport were measured using high-frequency instrumentation deployed at five stations set up at 10 m intervals along a transect oriented parallel to the wind direction from the top of the foreshore across the berm crest and upper beach (Figure 2A, B). Wind direction was from the NW at an angle of about 30° from shore parallel. The foreshore was quite gentle with a slope of 2° along the transect and a height difference of about 1.2 m between the berm crest and the low tide terrace (Figure 3B). Station 5 was set up near the upper foreshore, and station 4 was located on the berm crest. Station 3 was located on the gently sloping back surface of the berm crest and station 2 was within a slight dip in the topography that marked the junction between the landward-dipping berm slope and the seaward dipping beach surface on which station 1 was located.

Wind flow was measured using 2-D and 3-D sonic anemometers deployed on H-frames at stations 1 to 4 (Figure 2B). Each station consisted of two Gill 3-D Windmaster sonic anemometers positioned at 0.25 and 1.5 m above the bed, and a Gill 2-D windsonic anemometer mounted at a height of 3 m. These stations were positioned at 0, 10, 20 and 30 m along the profile (Figure 3B). An RM Young cup anemometer was positioned at station 5 near the top of the foreshore slope (40 m) at a height of 0.25 m. A mast with a vertical array of six cup anemometers was positioned near the back of the beach, but the data are not utilised in this paper.
Sand transport intensity was measured using two different sensor types: ‘Safire’ piezoelectric impact sensors (Baas, 2004; Barchyn and Hugenholtz, 2010) and Wenglor laser sensors (Davidson-Arnott et al., 2009; Hugenholtz and Barchyn, 2011). The Safires count only grains with sufficient momentum to register when they strike the 2 cm sensing ring, and grains that are too small, too slowly moving, or which only graze the sensor are not registered as counts. The sensitivity of the sensor also varies around the circumference of the instrument, particularly away from the two ‘sweet spots’ that mark the point of attachment of the leads to the sensing ring (Baas, 2004). It has been shown that the sensitivity varies from one sensor to another, and that sensitivity will vary over time particularly as the rubber sheathing weathers. The five Safires deployed here were selected after tests in the laboratory showed that their response to varying rates of sediment discharge in an air fall column was very similar. The instruments were deployed with their ‘sweet spot’ facing directly into the wind, thus removing concerns of directionality.

The laser sensor (Wenglor® model YH03PCT3) consists of a laser and photo sensor mounted within a U shaped frame with a spacing (path length) of 3 cm and a beam diameter of 0.6 mm. The instrument detects the drop in voltage at the photo sensor resulting from the passage of individual grains through the beam. The counting circuitry is contained within the instrument and is capable of detecting >2000 grains per second. Evaluation of the sensors indicates that they count all grains (or portions of grains) above the cut-off size that pass through the beam, irrespective of speed. As with all such sensors, they cannot distinguish
between particles entering the beam at exactly the same time so that a few particles are missed at moderate transport rates. At very large transport rates the actual count may go down (Hugenholtz and Barchyn, 2011) but this only occurs during very intense transport events. Transport rates here were never extreme enough for the latter to be a consideration (Davidson-Arnott et al., 2009).

Safires were deployed at all five stations but the sensor at station 4 failed early in the experiment. Wenglor laser sensors were deployed at stations 2 and 3 and also at a point half way between stations 1 and 2 (identified here as station 1.5). All sensors were sampled at 1 Hz over the course of the experiment. Because of capacity restrictions the data logger at station 5 had to be downloaded part way through the experiment resulting in a gap in data collection from the anemometer and Safire at that station. There was considerable noise in the data recorded by the anemometer at a height of 0.25 m at station 4 and, while it was possible to remove some of this through filtering, the uncertainties introduced into calculations of five-minute means meant that data from this sensor were discarded. Finally, sand transport rates were measured using vertical V traps (Nickling and McKenna Neuman, 1997) during four runs with a duration of 20 minutes (run 1) and 15 minutes (runs 2-4).

Moisture content and image processing

The long-term monitoring station at Greenwich consisted of three digital SLR cameras operated by a timer that triggered exposures every hour from May 2007 to May, 2010. The cameras were deployed on top of a 6 m mast located on
top of the 8 m high foredune crest (Figure 2C). The west camera had a field of view of approximately 100 m; the north camera covered a field of view of approximately 40 m alongshore; and the east camera covered an alongshore distance of 1.5 km. The east camera provided qualitative data on beach and weather conditions but the west and north cameras were used to extract numerical information such as moisture maps (MM) or fetch distances (Delgado-Fernandez et al., 2009). The short-term experiment described in this paper was conducted over a period of about six hours on October 21, 2007 within the field of view of the north camera (Figure 3A). Cameras were set to take exposures every five minutes during the short-term experiment.

The camera images were used to obtain time series of a number of key factors affecting aeolian sediment transport, such as moisture content and fetch distances (Delgado-Fernandez, 2011). In particular moisture maps (MM) were created by an automatic process (using Geomatica v9.1-Easi scripts) involving moisture calibration, image rectification and establishment of spatial extent and resolution. The process is described in detail by Delgado-Fernandez et al. (2009).

Moisture calibration consists of correlating surface brightness recorded by digital cameras with moisture content measured at the beach surface (Darke et al., 2009). Digital sampling of pixel brightness was conducted at beach locations where moisture samples were collected through surface scraping. Pixel brightness values were normalized against the brightness value of a ‘white board’ in order to compensate for the effect of different environmental light in the exposures due to sun angle and weather conditions. A calibration curve was
obtained by plotting normalized brightness against percent moisture content obtained from laboratory analysis, which allowed subsequent calculation of moisture values from pixel brightness values (Delgado-Fernandez et al., 2009). Image rectification transformed oblique imagery into georeferenced data by matching ground control points deployed within the field of view of the cameras with their corresponding map coordinates (surveyed with differential global positioning system). The raster cell size of the output rectified maps was 0.05 m and was set during the process of image rectification. This was the finest resolution possible given the area covered by an original pixel located in the farthest position (relative to the camera) subject to rectification. Shoreline position was digitized for each MM and vegetation cover was extracted using unsupervised classifications (Delgado-Fernandez et al., 2009). This allowed delimiting the beach width, which was used to establish the spatial extent of the MM.

Finally, near-surface moisture content was measured with a Delta-T Theta probe along the instrument transect at 2.5 m intervals. Sampling points were marked with rods (Figures 2B and 3) that were visible within the images taken by the camera. The probe was systematically used on the landward side of the rod, which allowed subsequent comparison between moisture values obtained with the Theta-probe and moisture values digitally sampled at the same locations in the corresponding MM (section 4). Moisture measurements were taken before runs 1-3 and after run 4 of the trap runs, at approximately 62, 112, 190 and 242 minutes from the start of the experiment. The length of the stainless steel rods of
the probe was shortened from the original 6 cm to 2 cm by using a 4 cm thick dielectric foam (Yang and Davidson-Arnott, 2005).

4. Results

Spatio-temporal variations in wind flow and sediment transport intensity

Wind speed averaged over five-minute periods are plotted in Figure 4A from the start of the experiment at 09:43 until 13:53 (250 minutes later) when recording ceased for the anemometers at stations 5, 3, 2 and 1. The general pattern recorded at all four stations was similar, with largest speeds recorded at station 5 near the top of the foreshore slope. There was an initial increase in mean speed to about 8.5 ms\(^{-1}\) at station 5 about 15 minutes after the start of recording followed by a decrease to about 7.3 ms\(^{-1}\). After 60 minutes the speed increased again, reaching a peak of approximately 9.5 ms\(^{-1}\) at about 90 minutes, remaining steady at this speed until 120 minutes and then decreasing to 8 ms\(^{-1}\) by 160 minutes. The corresponding averages for transport intensity are plotted for the four safires in Figure 4B and for the three Wenglor laser probes in Figure 4C. The temporal variations in mean sand transport intensity were somewhat more complex than for winds but showed a similar pattern with a peak early in the measurement period followed by a period with very low values for transport intensity and then much larger values between 90 and 120 minutes. This was followed by decreasing values until the end of the experiment, with the exception of the Wenglor at station 3, which showed more variability.
The spatial pattern of wind speed across the profile can be seen more clearly by comparing plots of five-minute means for five intervals (Figure 5B) that are representative of the range of conditions over the recording period, with the profile along the instrument line (Figure 5A). Except at the very end of the recording period the wind speed pattern showed largest speeds at station 5 near the top of the foreshore slope, slightly lower speeds at station 3 near the back of the berm crest and a further reduction in speed at station 2 located in the shallow dip of a remnant runnel at the landward end of the berm. Wind speed increased slightly at station 1 on the back beach. The pattern was consistent with acceleration due to flow compression as the wind traveled obliquely across the foreshore slope and then a small deceleration due to flow expansion in the lee of the berm crest. Unfortunately because of the problems with the lowermost sonic anemometer at station 4 it was not possible to determine if there was flow separation, but there is no consistent evidence for this in the records at station 3 and, because of the modest relief and relatively gentle apparent slope of the oblique transect, it seems likely that flow remained attached over this distance.

The largest transport intensity measured by the safires was generally found at station 3, landward of the berm crest, followed closely by the probe at station 5 (Figure 5C). Smaller values were recorded at stations 1 and 2; however, values for transport intensity at station 1 were almost always larger than those for station 2. The largest values for the Wenglor laser sensors were also recorded at station 3, with the smallest values recorded at station 2 and somewhat larger values at the probe located between stations 2 and 1 (station 1.5). This shows
that similar spatial patterns of sand transport intensity were recorded by these
two different sensor types.

Actual sand transport rates were measured using vertical integrating traps
that were deployed at all five stations. The trap data thus permit some evaluation
of transport conditions at station 4 compared to those at stations 5 and 3 (Figure
5D). Transport rates at station 4 on the berm crest were always larger than at
station 5. This pattern is to be expected if we assume that wind speed was likely
somewhat larger here because of flow acceleration due to flow compression up
the foreshore slope. Transport rates might also be expected to increase across
the profile with increasing fetch and decreasing surface moisture content.
Transport rates at station 4 were also larger than at station 3 except at the very
end of the experiment when wind speed had decreased considerably. At this time
it would be expected that the combination of light winds and large surface moisture
on the foreshore slope would result in conditions that were below the threshold
for much of the time at this location, and transport would be largely confined to
entrainment of dry sand from the surface of the berm and backshore.

Spatio-temporal variations in surface moisture

Images of the beach were acquired every five minutes, which resulted in a
total of 50 MM. An examination of the moisture conditions at this frequency
revealed that changes were difficult to discern, and hence, 30-minute time
intervals were used to assess changes in surface moisture (Figure 6A). A
representative sample corresponding to the middle period of the experiment is
shown here, from approximately 10:30 (55 min) to 13:00 (205 min). In addition to the moisture maps themselves, difference maps were generated by subtracting the data for each raster in one time period from the data for the same raster in the previous time period (Figure 6B). Three classes of surface moisture change were utilised: no change, drier, and wetter, with the last two classes corresponding to a decrease or increase in moisture content > 2% respectively.

The surface at the end of the first hour was slightly damp, ranging from 2 to 6% moisture content. A gradual increase in near-surface wind speed from 50 to 100 minutes (Figure 4A) led to erosion of a surficial layer of dry sand from the foreshore resulting in the west area beach surface becoming wetter at 85 minutes (Figure 6A). Wind speed continued to increase up to 115 minutes and remained steady and relatively strong until approximately 125 minutes. MM from 85 minutes to 145 minutes show a surface that progressively became drier on the backshore (due to deposition of dry sand) but especially wetter on the foreshore and west area (due to progressive erosion), as shown in the maps of changes in moisture content (Figure 6B). In contrast with the first part of the experiment, the foreshore became gradually dryer at 175 and especially at 205 minutes, with the formation of a continuous patch of dry sediment in front of all stations. This was due to a combination of sunnier conditions favouring evaporation across the beach surface and relatively gentle winds not capable of eroding the newly dried sand in the foreshore.

Moisture measurements with the Theta probe (Fig 7B) showed relatively small moisture content (5-7%) across the backshore, increasing moisture at the
berm and upper foreshore (6-8%), and the largest moisture content on the lower foreshore (9-12%). Moisture content was also re-sampled using the digital MM for the same locations at two grid resolutions: 0.05 m (Figure 7C) and 0.5 m (Figure 7D). The latter were created through a process of image coarsening, which combines cells into a larger cell and averages the values of the cells contained within it (Woodcock and Strahler, 1987). In the example used here, each 0.5 m cell represents the average of 100 0.05 m cells. The trends obtained from the MM followed, in general terms, those identified with the Theta probe. However the readings obtained from the coarser resolution maps (Figure 7D) showed simpler trend patterns, while moisture values sampled from the finest resolution maps (Figure 7C) showed unrealistically large variability. Average values from the 0.5 m renditions were therefore adopted for the purpose of comparison to the Theta probe values. Figure 8 shows a comparison of moisture content values obtained from the Theta probe along the profile versus those obtained from MM at 0.5 m grid resolution for different time intervals. In general, moisture values obtained from MM presented larger variability, with both drier and wetter zones than those obtained from the Theta probe. The remote sensing images seem highly sensitive to slight wetting or drying of the surface while this does not show up in the Theta probe results. Note that moisture readings taken close to the east side of the image (e.g., 15 m) at 112 minutes were greater than what would be expected from a dry surface. These are currently being explored as they could be a product of inaccuracies of moisture values at the edges of the
MM or a product of slight drizzle earlier in the morning, since they are no longer visible later in the day (see following section).

Co-evolution of surface moisture and sediment transport

Figure 9 shows time series for wind direction, wind speed, fetch distance, moisture content obtained from the 0.5 m grid resolution MM, and transport intensity given by the Safires for stations 1-5. Wind direction and speed data are from the 2-D sonic at a height of 3 m at station 4 and are used to provide general information of the incident wind field since this should be unaffected by growth of the boundary layer associated with compression of flow up the foreshore slope. Gaps in the moisture and fetch graphs correspond to periods of time when no images were taken. Wind direction was initially very constant, from approximately 10° to the E of the instrument transect (aligned with a wind direction of 180°), gradually becoming more onshore (where directly onshore is 270°) and staying at 20° to the E of the instrument transect by the end of the experiment. Figure 3C shows the approximate location of the shoreline at the beginning and end of the experiment, with small changes at this particular location. However, the combined effect of a rising tide and gradual change in wind direction resulted in a decrease of available fetch distance towards the end of the experiment, especially in the landward stations because of differences in beach configuration (see DEM of Figure 3A). The fetch distance decreased over the course of the experiment by 9 m at station 5 and 22 m at station 1 (Figure 9). This reduced fetch has significant implications for sediment transport potential across the
beach and may explain, in part, why transport intensity after about 140 minutes was greatly reduced (Fig 9E).

The beginning of the experiment was marked by cloudy conditions and occasionally light drizzle, which when combined with relatively calm wind conditions may have played a role in the reduction in sand transport during the first hour. The drizzle ended about 50 minutes into the experiment, and soon after that the clouds cleared so that evaporation due to solar radiation became a factor in the reduction of surface moisture content during the latter half of the experiment. Increased sediment transport during the second hour likely reflects both increasing wind speed and surface evaporation. Note that while this is the trend at stations 4 and 3 (Figure 9D), station 5 (closer to the foreshore) was characterised by increasing surface moisture content due to progressive stripping of the dry surface sediment and exposure of the wetter materials below.

The foreshore and lower beach were erosion zones whereas the back beach was a deposition zone. At the end of the transport event, measurements showed that about 5 cm of scour had occurred at station 5 with a smaller amount at station 4 on the berm crest (Figure 10A). In contrast about 5 cm of deposition was measured at station 3, 3 cm at station 2, and about 1 cm at station 1 (Figure 10B). The configuration of this spatial zonation from shoreline to the foredune toe became gradually aligned normal to the incoming wind direction as the event progressed, with a transition zone in between strongly affected by beach topography and shoreline position. The transition zone was a relatively wide area (roughly around station 4) with a mixture of moist and dry conditions, and
characterized by a fairly corrugated edge line resulting from transport processes acting on a very complex surface due to the presence of stones, flotsam, and footprints characterising the erosional surface (Figures 6 and 10C). This level of complexity favours spatial approaches such as MM and remotely sensed data over spot measurements along single transects such as those collected from Theta-T probes, which do not provide enough detail on spatial textures.

5. Discussion

Comparison of Theta-T and MM moisture values raise a number of questions regarding the assessment of surface moisture content as a primary control on aeolian sediment transport. Values from MM of 0.5 m resolution showed a better correlation with the Theta-T probe values, which are reliable indicators of surface moisture content when there are smooth transitions in moisture gradients along a given beach transect. Values from MM at 0.05 m resolution contain a large degree of variability in moisture readings, and the question is whether such variation is ‘noise’ or a true representation of the nature of the evolving beach surface during a transport event. When transport was active, there were discrete zones of dry sand that were being deposited beside a moist transport surface that was recently eroded. For example, Figure 2B shows a complex surface with moisture patches alongside dry sand as a result of active aeolian streamers and deposition/erosion processes. In this instance, the MM appear to be yielding ‘noisy’ but detailed information that may be representative of the complex evolution of the beach surface through time. Further analysis is
being undertaken to investigate these micro-scale complexities and to assess the accuracy of moisture data at the edges of MM.

A related issue has to do with quantifying a characteristic moisture value that should be used in explanations of sediment transport as a function of wind, moisture, and fetch distances. The 0.05 m resolution MM provide detailed representations of beach surface moisture but sediment transport rates might correlate better with averaged moisture values over a given area. Furthermore, transport rates could be better explained using some measure of the spatial distribution of moisture values over a set distance (e.g., 1 m, 2 m, or 10 m) in front of a given trap. MM provide an excellent platform for these types of investigations because they allow analysis of moisture content at different spatial resolutions.

In addition to providing a measure of the actual moisture content of the upper-most sand grain layer, in contrast to the depth-averaged sample (from 1.5-6 cm thickness) from the Theta probe (Atherton et al., 2001; Yang and Davidson-Arnott, 2005; Schmutz and Namikas, 2011), the process of cell aggregation of fine resolution (0.05 m) MM into coarser resolution (0.5 m) MM offers the advantage of providing more representative spatially-average values. Detailed sampling of surfaces with moderate moisture content using the Theta probe have shown that there can be a large degree of variability within an area as small as 1 m² (Yang and Davidson-Arnott, 2005; Edwards and Namikas, 2009). While Theta probes provide point measurements that are assumed to be representative of a given area, a cell from the coarse resolution MM presented here would represent
the average of 100 moisture measurements over an area of 0.5 m². Taking a similar size sample with a Theta probe would be intrusive and obviously highly time consuming.

While surface moisture is an important control on aeolian transport (Belly, 1964; McKenna Neuman and Nickling, 1989; Namikas and Sherman, 1995; Cornelis and Gabriels, 2003; Wiggs et al., 2004b; Davidson-Arnott et al., 2008) results from this study demonstrate that transport is in turn an important control on surface moisture content as the beach surface evolves through localized erosion and deposition. Surface moisture is related to transport in two opposing ways. Safires at stations 5 (foreshore) and 3 (back beach) recorded a peak of transport intensity from approximately 80 to 120 minutes, coinciding with the increase of wind speed. However, surficial moisture content increased at station 5 but decreased at station 3 (Figure 9). The combination of flow acceleration up the foreshore slope and a short fetch (no sediment source upwind) transformed the foreshore (station 5) into an erosive zone. Hence, relatively large values for moisture content at this location were the result of sediment erosion and landward transport. This is in line with observations on sediment transport dynamics on a drying beach by Nield et al. (2011) who reported erosion of dry sediment leading to an increase of surface moisture due to the exposure of the underlying, wetter surface. Large transport rates at station 3 (back beach) were associated with significant accumulation of dry sand and hence small values of surface moisture content. Details of the temporal evolution of moisture content in these two different zones reveal complex feedbacks between moisture, wind
speed, sand transport, and fetch distances where moisture is not only a control for sediment movement but also a result of patterns of erosion and accumulation. The largest transport rates at stations 5 and 3 were recorded during the first trap run (96 minutes; Figure 5D), coinciding with the strongest wind speeds and the peak of saltation recorded by the Safires (Figure 9). As suggested by the MM (Figure 6A) most of the sand collected during this run at station 5 was the result of the erosion of a relatively dry layer from the foreshore, which in turn increased the surface moisture content there. Transport rates at station 5 during the second run (127 minutes) were smaller than at station 1 (Figure 5D) despite strong winds of 9-9.5 ms\(^{-1}\) similar to those recorded during run 1. Hence the erosion of the layer of dry sediment from this zone during run 1 produced larger levels of surface moisture during run 2, which in turn limited further erosion. At stations 3-1, dry sediment stripped from the foreshore was deposited progressively leading to drier surface sediment. These deposits are easily mobilized under stronger wind conditions, but boundary layer evolution leads to reduced surface shear stress and hence, deposition toward the toe of the dune.

6. Conclusion

This paper demonstrates the applicability of a remote sensing system to measuring beach surface moisture over a short-term experiment. The detailed temporal and spatial resolution of the evolution of surface moisture provided by the time-lapse photography provides an improved explanation for the spatial and temporal patterns of sand transport, and the changes in the beach surface
resulting from scour and deposition. The conclusions can be summarised as follows:

1) Spatial patterns of erosion, transport and deposition over the foreshore and berm reflect the changing boundary layer produced by flow compression up the foreshore slope and then expansion over the reverse slope on the landward side of the berm;

2) Temporal patterns of sand transport intensity reflect both changes in wind speed and direction (fetch distance) over the event, upwind sources of available dry sediment, and the effects of changes in surface moisture;

3) Fine resolution moisture maps offer the possibility of data aggregation through pixel coarsening and may be used to explore appropriate moisture values for different applications. While averaged moisture values obtained from coarser resolution (0.5 m) MM were more similar to those obtained with the Theta probe these may be smoothing out important micro-scale information. Further research should be conducted to determine the appropriate temporal and spatial resolution of moisture measurements.

4) The high resolution of the surface MM permits the identification of changes in moisture over the whole surface that reflect external conditions such as solar radiation input and precipitation as well as internal conditions produced by the erosion (wetting) and deposition (drying) of sand.

5) The remote sensing system thus provides a relatively simple approach to measuring the evolution of the beach surface during a transport event, and
to providing more comprehensive explanations of processes of erosion
and transport and changes in surface conditions.

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Figure 1: Location of study site at Greenwich Dunes, Prince Edward Island (Canada).
Figure 2: A) General view of study site and instruments; B) close-up of instruments setup. The transect consisted of four stations, each with one 2D and two 3D sonic anemometers located at 3, 1.5, and 0.25 m over the beach surface, and one station with a cup anemometer at 0.25 m over the beach surface. Traps and safires were co-located at each of the stations and moisture points selected every 2.5 m. The transect ran into the incoming wind direction (oblique onshore) and ended in a tower of cup anemometers close to station 1; C) close-up of camera system.
Figure 3: A) Digital elevation model (DEM) georectified from 1 m spot resolution differential global positioning system (DGPS) data. Topographic data was gathered on the beach surface and up to the dune toe and thus the DEM does not include the dune. The position of each station is marked with the station number. Stations 3-5 and a total of 15 moisture points were within the north camera field of view (inserted as a semi-transparent image); B) topographic profile along the instruments transect, showing the approximate location of the shoreline position at the beginning (shoreline start) and end of the experiment.
Figure 4: Temporal variations in wind speed and sediment transport intensity. A) five-minute averages for wind speed (ms\(^{-1}\)) measured at stations 1, 2, 3 and 5 at 0.25 m height; B) average saltation intensity (counts s\(^{-1}\)) measured by Safire probes at the same stations; C) average saltation intensity (counts s\(^{-1}\)) measured by Wenglor laser sensors at stations 2, 3 and midway between 1 and 2. Note that the grain counts measured by the Wenglors are much larger than those measured by the Safires (see Davidson-Arnott et al., 2009).
Figure 5: Spatial pattern of winds and sediment transport. A) beach profile along instruments transect; B) mean wind speeds for 6 five minute periods; C) transport intensity measured by the Safire probes for the same 5 minute periods as B); D) transport rate measured by vertical traps. Run 1 had a duration of 20 minutes and the other three runs a duration of 15 minutes. Times are for the end of the run.
Figure 6: Temporal variations in moisture content. A) temporal sequence of moisture maps every 30 minutes; B) maps showing the change in moisture content relative to the previous moisture map. Only relative changes over +/-2% have been considered, with drier areas (change > 2%) shown in orange and wetter areas (changes > -2%) in green. The black dot shows the location of station 4.
Figure 7: A) Beach profile along the instruments transect; B) Measurements of near surface moisture (upper 2 cm) taken at four times along the instrument profile using the Delta-T Theta Probe; C) Surface moisture at the same four times and locations (within the field of view of the camera) estimated from moisture maps at the original grid resolution of 0.05 m; D) Measurements of surface
moisture from moisture maps at a coarser resolution of 0.5 m.

Figure 8: Plot of moisture values from the Delta-T Theta probe and moisture maps (MM) at 0.5 m grid resolution at three different time intervals. In general terms, the MM values show a larger variability of moisture values and sharper differences between the wettest and driest ends.
Figure 9: Five minute averages for incident wind direction (A) and wind speed (B) recorded at 3 m high over the beach surface at station 4, fetch distances (C), surface moisture content for stations 3-5 (within the field of view of the camera; D), and sediment transport intensity (E). Gaps in the data correspond to periods of time when no images were taken.
Figure 10: Photographs of the beach surface at the end of the transport event: A) close-up of station 4; B) ripples formed in the dry sand accumulation at station 2; C) view of station 4 looking landward showing the eroded and relatively damp surface with a transition to deposition towards station 3 in the background.